

Sea Ice Mechanics

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With thanks to

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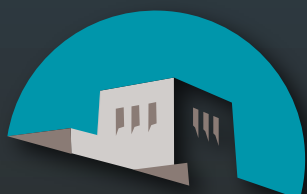
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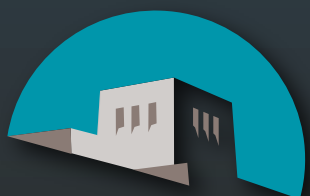
Research Focus

New Sea-Ice Constitutive Model

- Elastic-Decohesive Model

Numerical Method

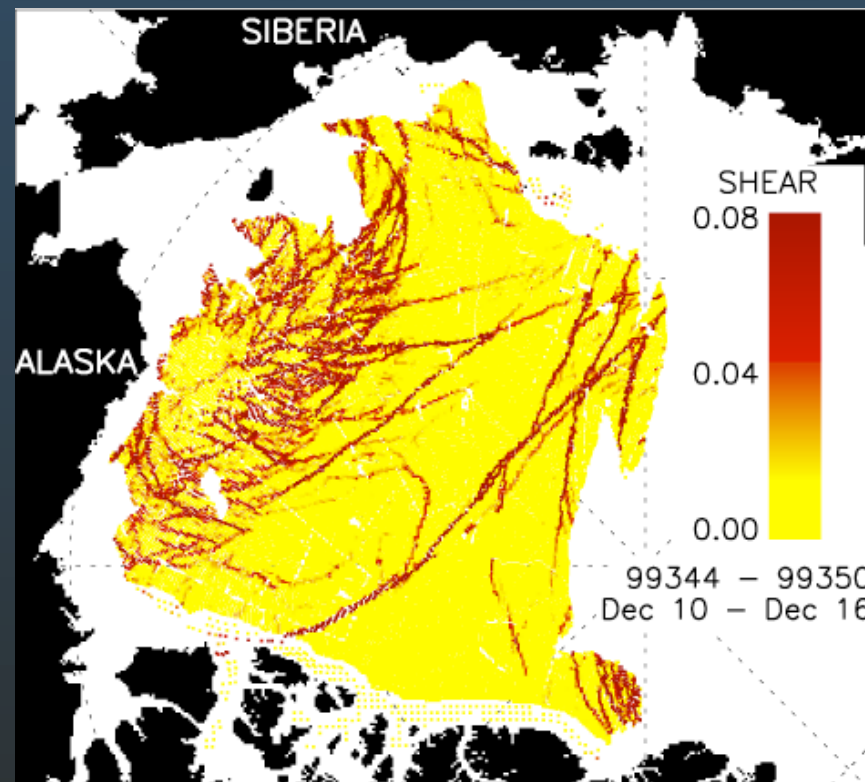
- Material-Point Method



Why A New Ice Model?

The viscous-plastic model is an isotropic model based on a 100 km scale in which it was assumed that cracks, ridges and leads were randomly distributed.

RGPS analysis of satellite images shows large ice deformation events occurring in long-lasting linear features that appear to correspond to displacement (or velocity) discontinuities in the deformation field due to leads.



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Elastic-Decohesive Sea Ice Model

Overall Objective: Numerically simulate “linear kinematic features” (eg. leads and ridges)

Initial Focus: Prediction and appearance of leads

Proposed Approach: Elastic-Decohesive Model

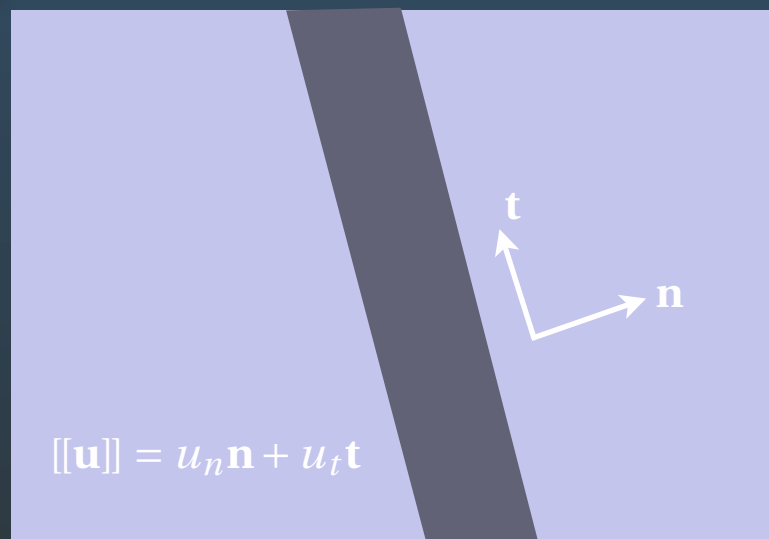
For thick first-year ice and multi-year ice, we assume most deformation occurs due to discontinuities in the displacement field.

Ice is quasibrittle so we can borrow from models of concrete and rock.



Elastic-Decohesive Model

- Intact ice modeled as elastic
- Leads modeled as discontinuities
- Model predicts initiation of a lead and its orientation
- Traction is reduced with lead opening until a complete fracture forms
- Implementation similar to elastic-plastic model

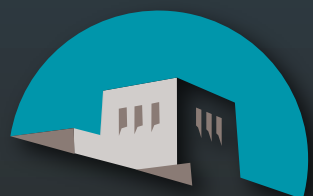


Schreyer, H., L. Monday, D. Sulsky, M. Coon, R. Kwok (2006), Elastic-decohesive Constitutive Model for Sea Ice, J. of Geophys. Res., 111, C11S26, doi:10.1029/2005JC003334.

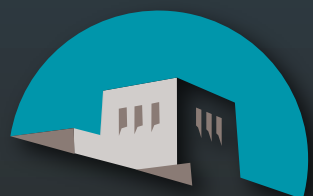
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Simulate a Region of the Beaufort

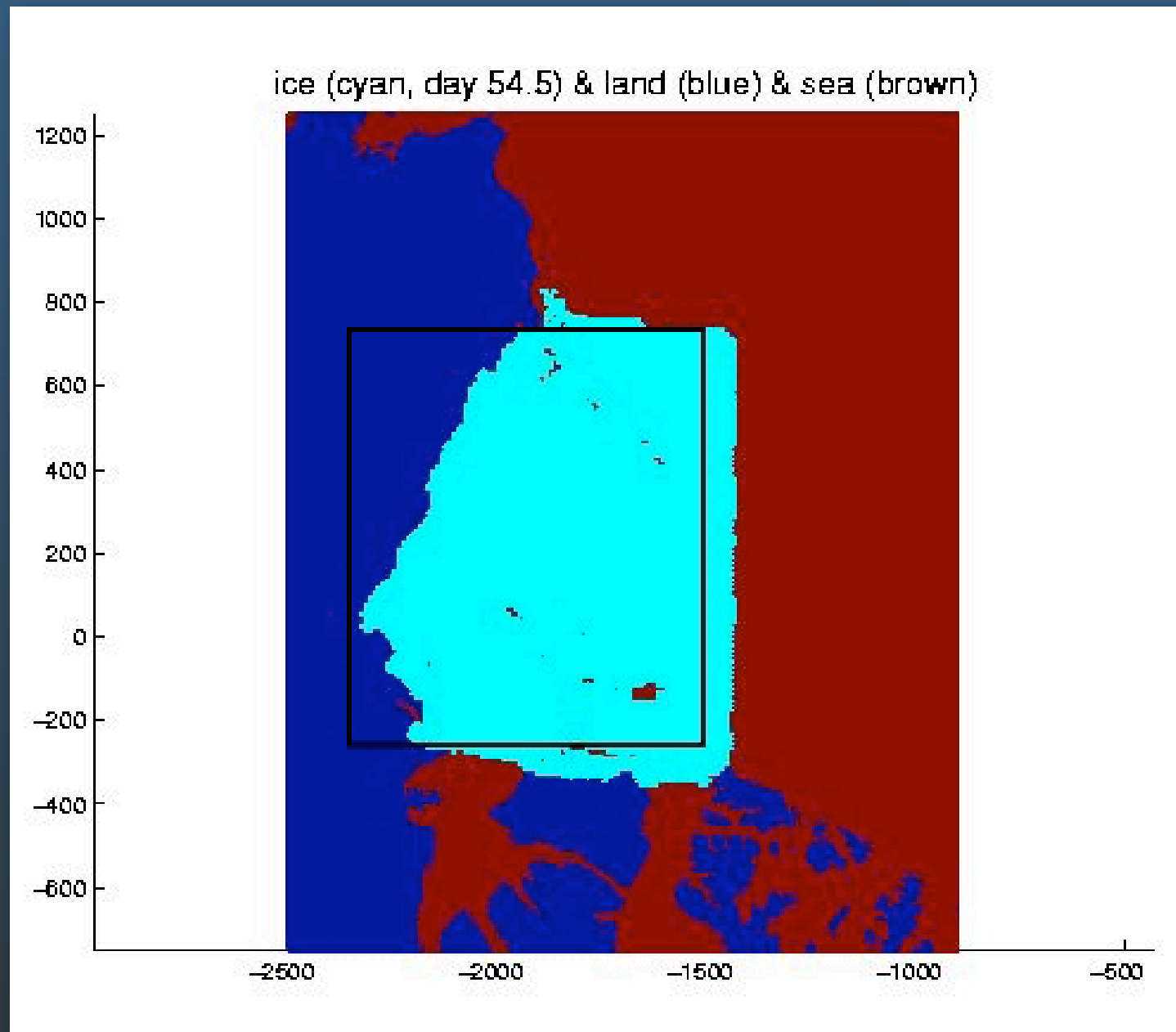


Simulate a Region of the Beaufort



Problem Set Up

Simulate 16 days in Feb/Mar, 2004



$$E = 1 \text{ MPa}$$

$$\nu = 0.36$$

$$\tau_{nf} = 25 \text{ KPa}$$

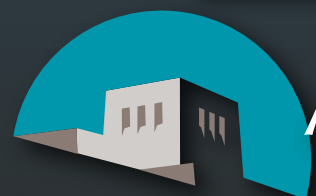
$$\tau_{sf} = 15 \text{ KPa}$$

$$f'_c = 125 \text{ KPa}$$

$$u_0 = 400 \text{ m}$$

$$S_m = 4$$

Apply known winds and ocean currents from a model.



Material-Point Method

Solves standard momentum equation

$$(\rho h) \frac{d\mathbf{v}}{dt} - \mathbf{t}_a - \mathbf{t}_w + (\rho h) f_c (\mathbf{e}_3 \times \mathbf{v}) - \nabla \cdot (\boldsymbol{\sigma} h) = 0$$

—————

Inertia

—————

Coriolis

—————

Stress div

Air Drag:

$$\mathbf{t}_a = c_a \rho_a \|\mathbf{v}_a\| \mathbf{R}_a \mathbf{v}_a \quad \mathbf{R}_a = \begin{bmatrix} \cos \alpha & -\sin \alpha \\ \sin \alpha & \cos \alpha \end{bmatrix}$$

Water Drag:

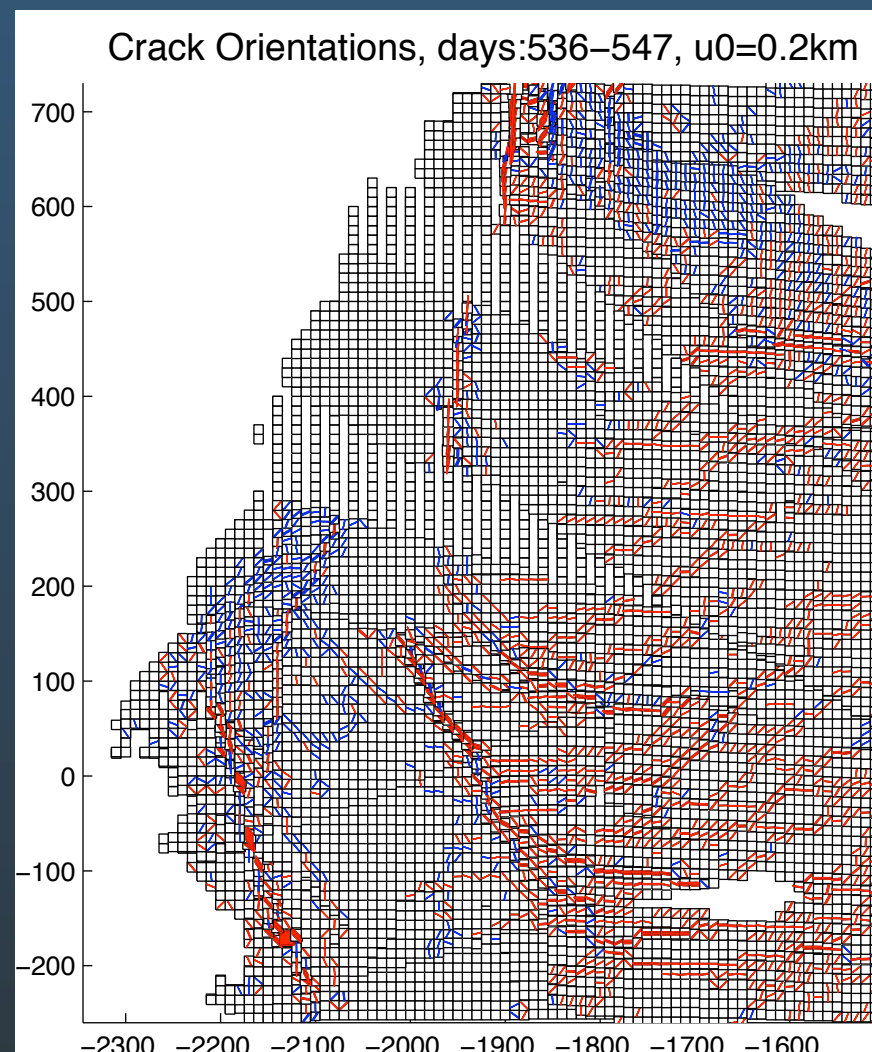
$$\mathbf{t}_w = c_w \rho_w \|\mathbf{v} - \mathbf{v}_w\| \mathbf{R}_w (\mathbf{v} - \mathbf{v}_w) \quad \mathbf{R}_w = \begin{bmatrix} \cos \beta & -\sin \beta \\ \sin \beta & \cos \beta \end{bmatrix}$$

Plus ice thickness distribution and Bitz-Lipscomb thermodynamics



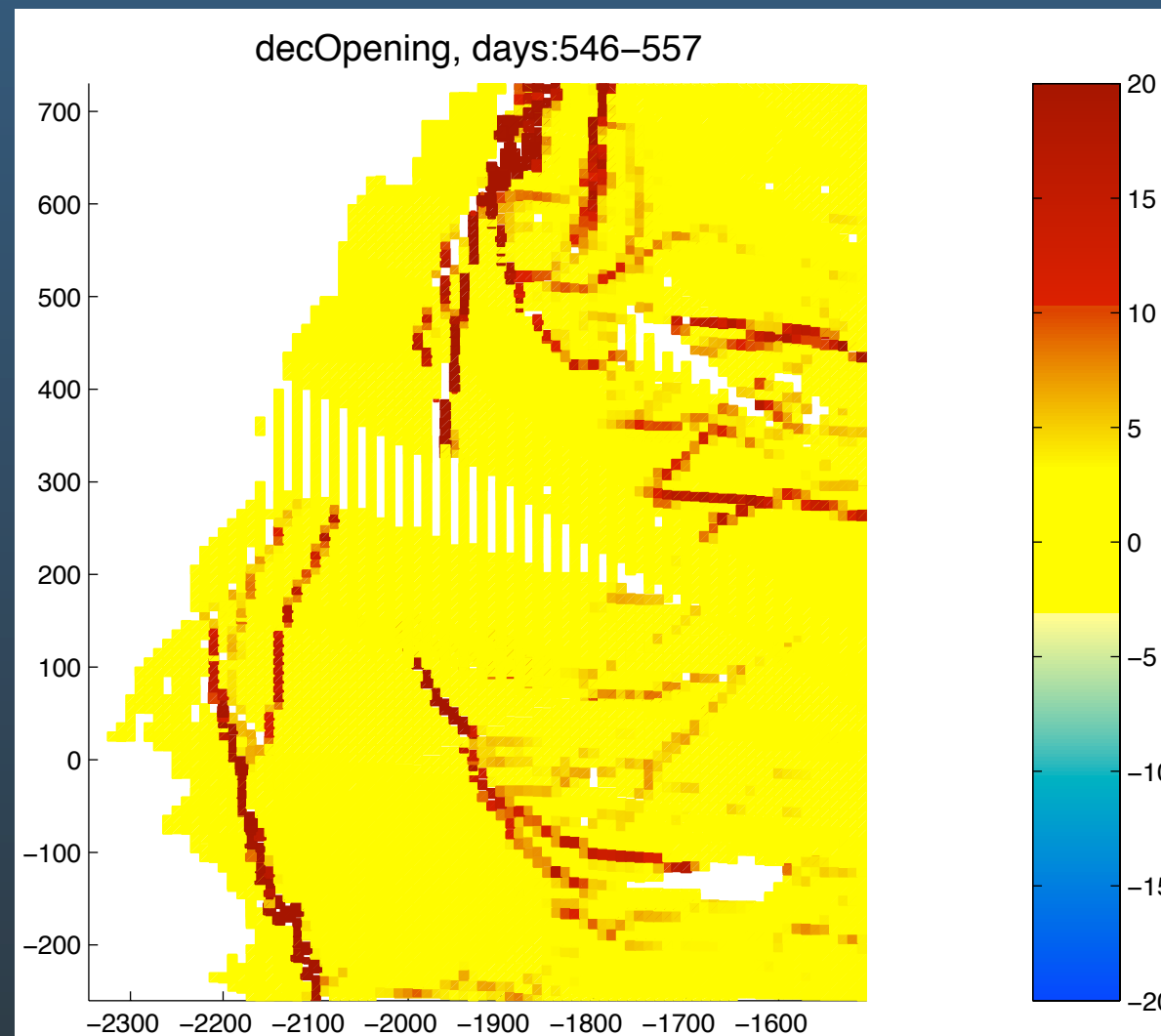
Initialization

Where are existing leads?

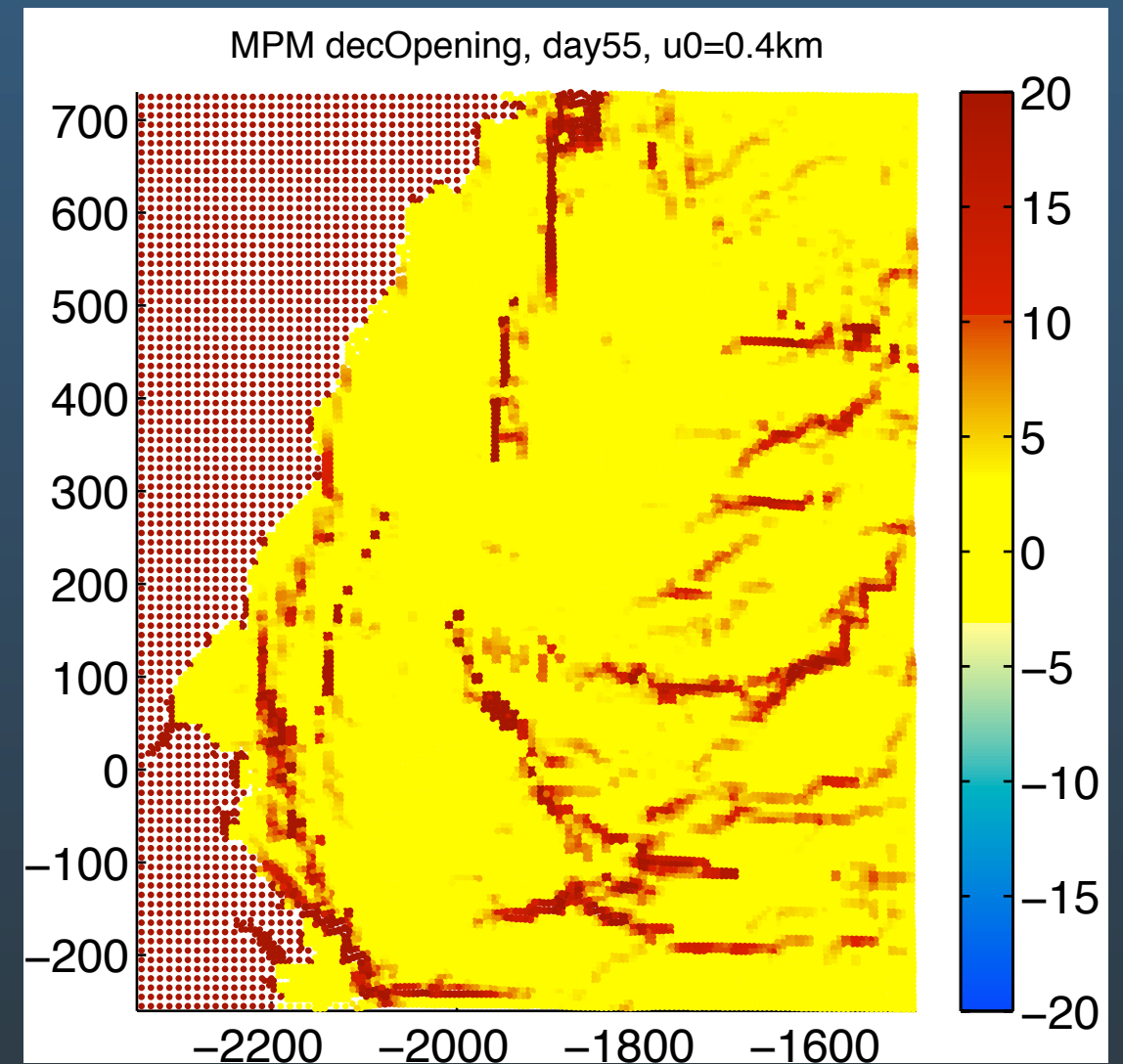


Fracture Patterns in the Beaufort

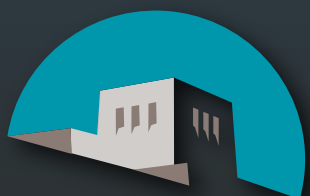
day 55



Observation

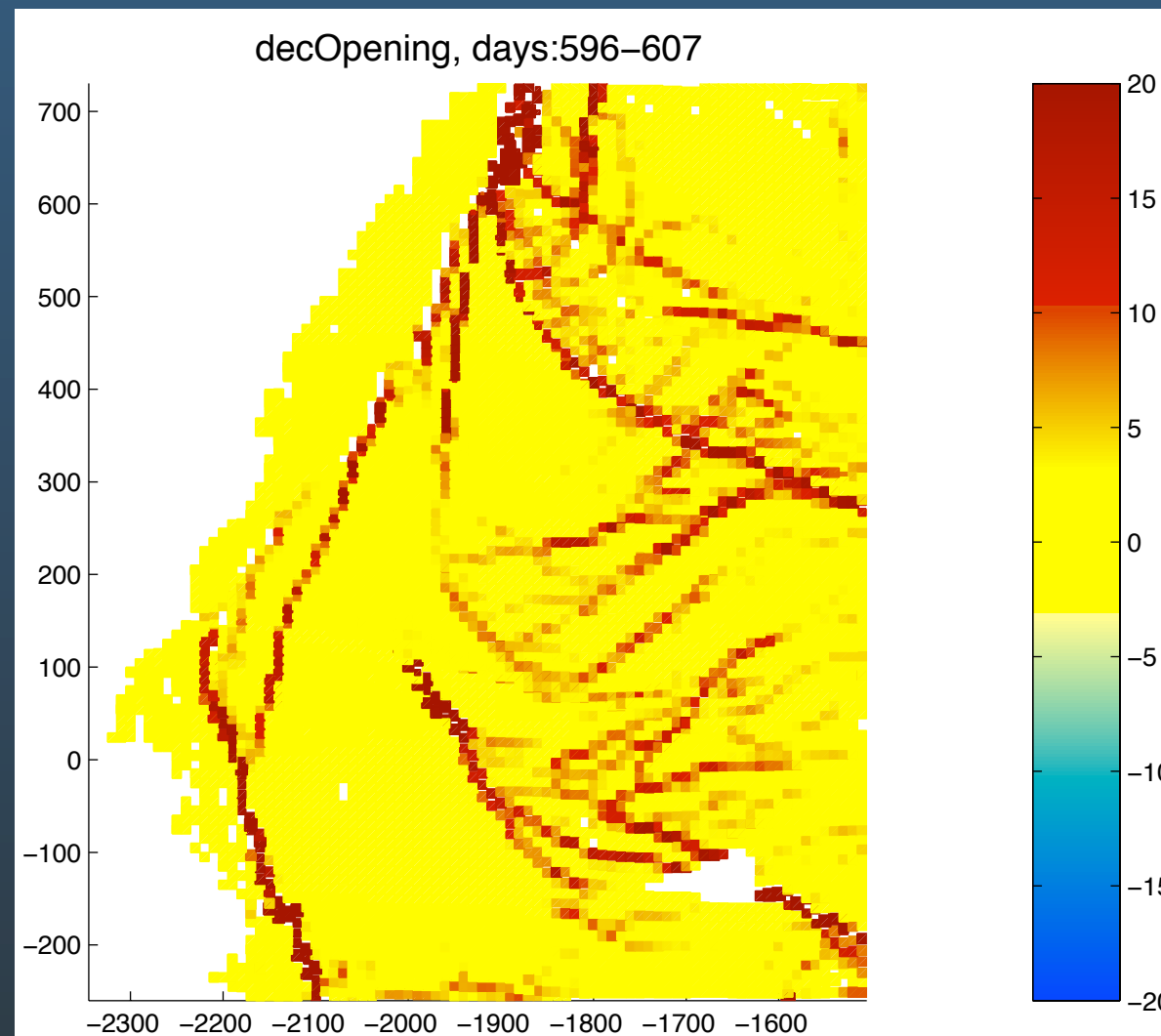


Simulation

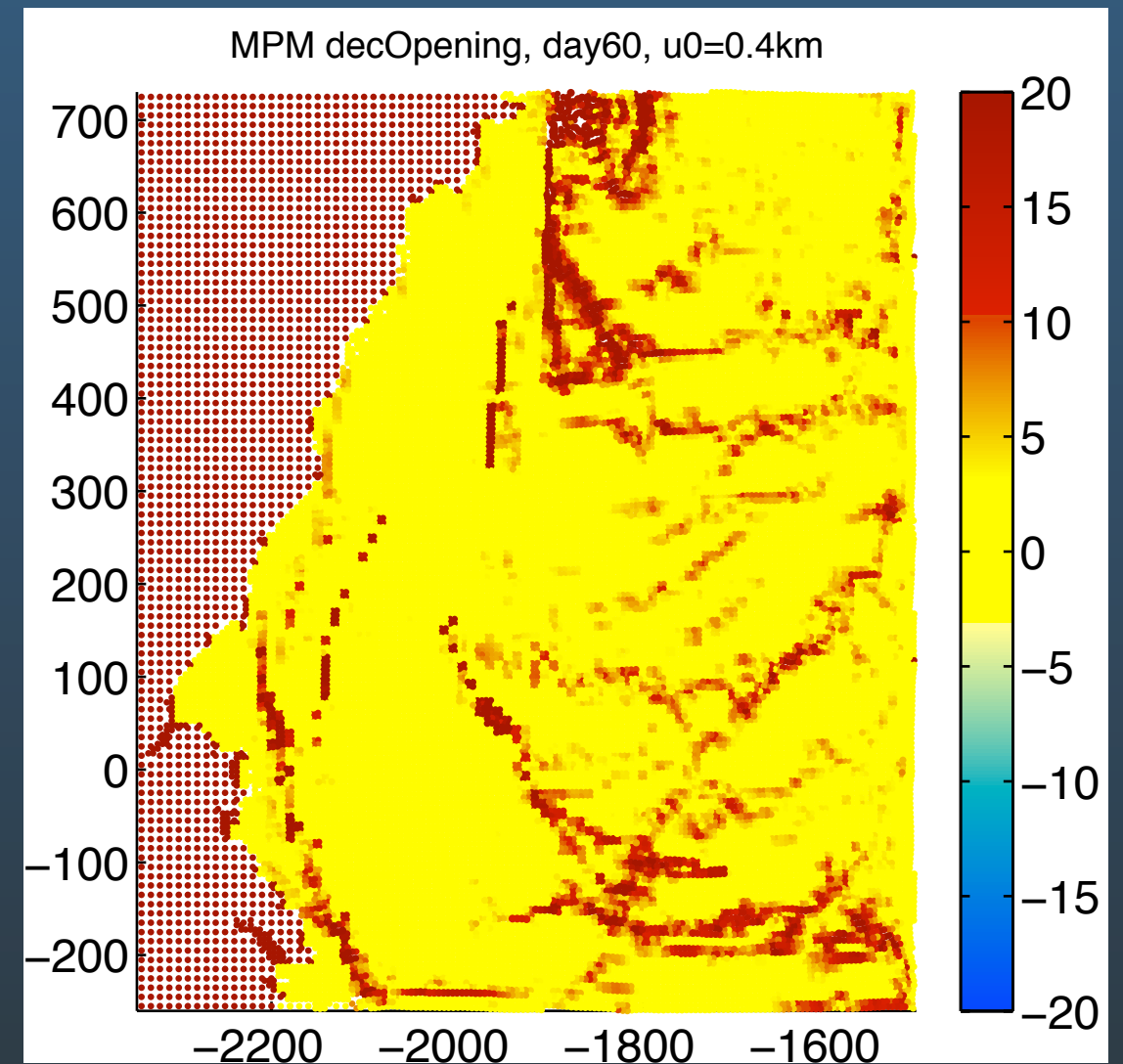


Fracture Patterns in the Beaufort

day 60



Observation

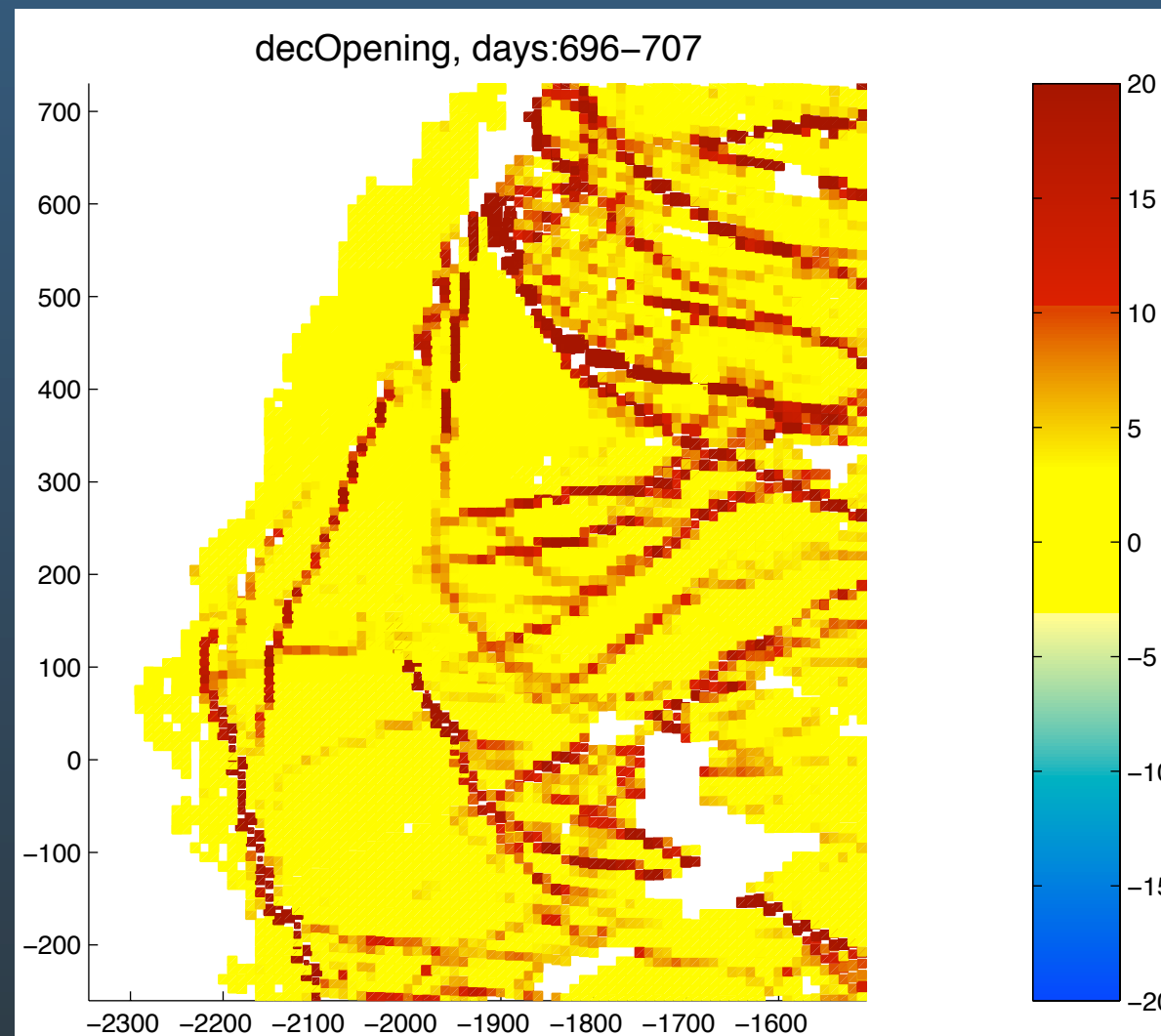


Simulation

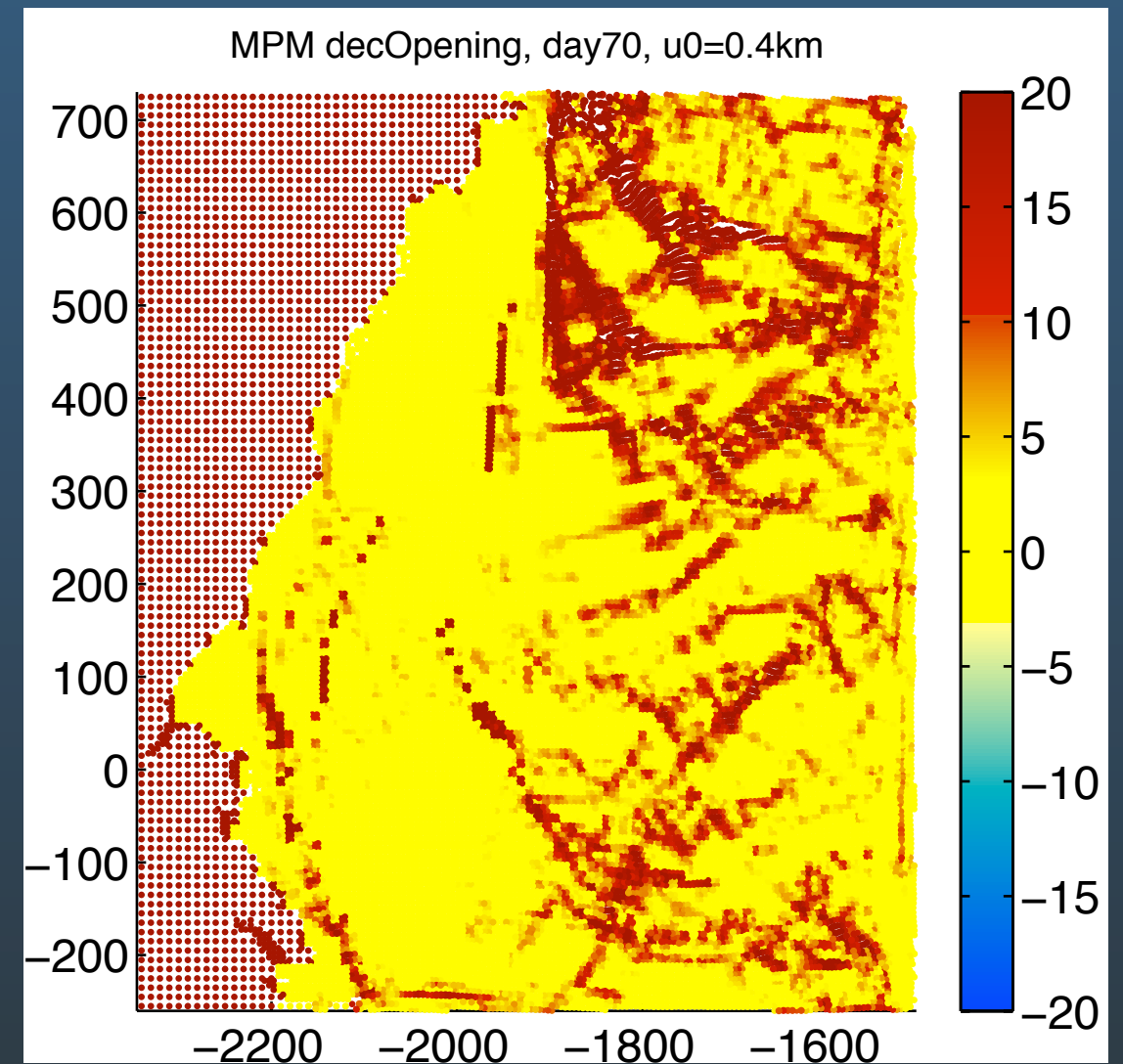


Fracture Patterns in the Beaufort

day 70



Observation



Simulation



Notes

Calculations were done

- without tuning parameters
- with crude initial guess
- with no refreezing of leads
- with errors from preprocessing satellite data



Summary

Features of elastic-decohesive model:

- Stress state at which leads initiate
- Orientation of lead at initiation
- Evolution of lead (softening)
- Existing material weakness
- Implemented in plasticity framework

Work in progress:

- Initial conditions
- Freezing model
- Coupling to ocean
- Metrics

